## Transition to geostrophic convection: the role of the boundary conditions

Ruddie P.j. Kunnen<sup>1</sup>, Rodolfo Ostilla-Mònico<sup>2</sup>, Erwin P. Van Der Poel<sup>2</sup>, Roberto Verzicco<sup>\*†3,2</sup>, and Detlef Lohse<sup>2</sup>

<sup>1</sup>Fluid Dynamics Laboratory – Department of Applied Physics and J.M. Burgers Centre for Fluid Dynamics, Eindhoven University of Technology, P.O. Box 513, 5600 MB Eindhoven, Netherlands <sup>2</sup>Physics of Fluids Group – Mesa+ Institute and J.M. Burgers Centre for Fluid Dynamics, University of

Twente, P.O. Box 217, 7500 AE Enschede, Netherlands

<sup>3</sup>Dipartimento di Ingegneria Industriale – University of Rome "Tor Vergata", Via del Politecnico 1, Roma 00133, Italy

## Abstract

Rotating Rayleigh–Bénard convection, the flow in a rotating fluid layer heated from below and cooled from above, is used to analyse the transition to the geostrophic regime of thermal convection. In the geostrophic regime, which is of direct relevance to most geo- and astrophysical flows, the system is strongly rotated while maintaining a sufficiently large thermal driving to generate turbulence. We directly simulate the Navier–Stokes equations for two values of the thermal forcing, i.e.  $Ra = 10^{10}$  and  $Ra = 5 \cdot 10^{10}$ , a constant Prandtl number Pr = 1, and vary the Ekman number in the range  $Ek = 1.3 \cdot 10^{-7}$  to  $Ek = 2 \cdot 10^{-6}$  which satisfies both requirements of super-criticality and strong rotation. We focus on the differences between the application of no-slip vs. stress-free boundary conditions on the horizontal plates. The transition is found at roughly the same parameter values for both boundary conditions, i.e. at  $Ek \approx 9 \times 10^7$  for  $Ra = 1 \times 10^{10}$  and at  $Ek \approx 3 \times 10^7$  for  $Ra = 5 \times 10^{10}$ . However, the transition is gradual and it does not exactly coincide in Ek for different flow indicators. In particular, we report the characteristics of the transitions in the heat transfer scaling laws, the boundary-layer thicknesses, the bulk/boundary-layer distribution of dissipations and the mean temperature gradient in the bulk. The flow phenomenology in the geostrophic regime evolves differently for no-slip and stress-free plates. For stress-free conditions the formation of a large-scale barotropic vortex with associated inverse energy cascade is apparent. For no-slip plates, a turbulent state without large-scale coherent structures is found; the absence of large-scale structure formation is reflected in the energy transfer in the sense that the inverse cascade, present for stress-free boundary conditions, vanishes.

## References

- Chandrasekhar, S. 1961 Hydrodynamic and Hydromagnetic Stability. Oxford: Oxford University Press.
- [2] Ecke, R. E. 2015 Scaling of heat transport near onset in rapidly rotating convection. Phys. Lett. A 379, 2221–2223.

\*Speaker

<sup>&</sup>lt;sup>†</sup>Corresponding author: verzicco@uniroma2.it

- [3] Julien, K., Knobloch, E., Rubio, A. M. Vasil, G. M. 2012 Heat transport in low-Rossbynumber Rayleigh–Bénard convection. Phys. Rev. Lett. 109, 254503.
- [4] Kunnen et al 2016 Transition to geostrophic convection: the role of the boundary conditions . J. Fluid Mech. (to appear)